

Laplace transform analytic element method for transient porous media flow

Kristopher L. Kuhlman and Shlomo P. Neuman

Department of Hydrology and Water Resources

University of Arizona, Tucson, AZ 85721

Abstract. A unified theory of the Laplace transform analytic element method (LT-AEM) for solving transient porous media flow problems is presented. LT-AEM applies the analytic element method (AEM) to the modified Helmholtz equation, the Laplace-transformed diffusion equation. LT-AEM uses superposition and boundary collocation with Laplace-space convolution to compute flexible semi-analytic solutions from a small collection of fundamental elements. The elements discussed are derived using eigenfunction expansions of element shapes in their natural coordinates. A new formulation for a constant strength line source is presented in terms of elliptical coordinates and complex-parameter Mathieu functions. Examples are given illustrating how leaky and damped wave hydrologic problems can be solved with little modification using existing LT-AEM techniques.

Keywords: analytic element, Laplace transform, diffusion equation, modified Helmholtz equation, elliptical coordinates, transient line source, Mathieu functions

Abbreviations: AEM – analytic element method; LT-AEM – Laplace transform analytic element method

1. Introduction

The analytic element method (AEM) provides semi-analytic solutions to linear porous media flow problems through superposition of fundamental solutions. The original development of AEM is due to Strack and his co-workers at the University of Minnesota [1]. During its initial development, it was compared to the boundary element method [2], but the eigenfunction expansion approach discussed here may be considered a special case of the spectral method (see [3, Sect. 3.1] and [4, App. C]). Each AEM element satisfies the governing equation, while spectral elements typically do not. The majority of AEM applications to date have been concerned with vertically-averaged steady-state groundwater flow (the two-dimensional (2D) Laplace and Poisson equations in the horizontal plane). AEM has been extended to three-dimensional [5], transient [6], multi-layer [7], and linearized unsaturated [8] flow problems. Review papers [9–11] and textbooks by Strack [12] and Haitjema [13] cover the fundamentals, discuss applications, and mention some recent advances.

AEM partially fills a gap between analytic solutions derived for simple geometries (e.g., radially-symmetric flow to a well) and distributed-parameter gridded models (e.g., finite element methods). AEM is well-suited for time-independent boundary value problems; applications to transient diffusion have proceeded in several directions.

The earliest extension of AEM to transient flow [14] was discontinuous in time, using a grid to simulate transient storage. The corresponding space discretization offset the mesh-free benefit normally associated with AEM. Another early approach [15] combined steady and transient elements, using line and area sources to model transient effects; transient storage effects were assumed piecewise-constant in time and the method required zero net withdrawal of water from the aquifer. While not an AEM solution, Butler and Liu [16] developed a solution for transient flow to a well in the presence of a single circular inhomogeneity, using an approach similar to that taken here. Bakker [17] used a temporal Fourier transform to apply the AEM to problems comprised of a finite number of temporal harmonics. Strack [18] described a general AEM approach in which localized transient perturbation elements are superimposed on a confined steady background that uses finite differences in time.

Furman and Neuman [6] first used AEM to solve the Laplace-transformed diffusion equation. LT-AEM numerically back-transforms the solution into the time domain using an inverse



47 Laplace transform algorithm. In contrast to the Fourier transform approach, the use of the
 48 Laplace transform obviates the need for periodicity and can incorporate initial conditions.
 49 We illustrate LT-AEM elements constructed using eigenfunction expansion and Laplace-space
 50 convolution. Steady-state multi-layer aquifer systems [7] and linearized steady unsaturated flow
 51 [19, 20] also lead to the Helmholtz equation. We show how these type of homogeneous distributed
 52 sources can be incorporated into the LT-AEM. The elements outlined here are restricted to
 53 simple geometries (i.e., circles and ellipses), but other techniques (e.g., those utilized in spectral
 54 element modeling [3, Chap. 17]) can be used to extend LT-AEM to more general geometries.

55 2. Laplace Transform AEM

56 Hydraulic head in a transient, 2D, confined, elastic aquifer is described by the diffusion equation,

$$Kb\nabla^2 h(\mathbf{x}, t) + bG = bS_s \frac{\partial h(\mathbf{x}, t)}{\partial t}; \quad (1)$$

57 where $h(\mathbf{x}, t)$ is vertically-averaged hydraulic head [L], b is aquifer thickness [L], G is a volumetric
 58 source term [$1/T$], K is the hydraulic conductivity [L/T], and S_s is specific storage [$1/L$]. K
 59 is assumed isotropic; both K and S_s are assumed homogeneous. For horizontal 2D flow, a unit
 60 aquifer thickness is assumed for simplicity (unless stated otherwise), without loss of generality.
 61 2D vertical-plane flow could also be simulated with this approach, e.g., flow under a wide dam
 62 on a permeable foundation.

63 In AEM it is standard to work with discharge potential [L^3/T], $\Phi = bKh + C$, where C is an
 64 arbitrary reference [12] that we conveniently set to zero. Applying the Laplace transform [21,
 65 Chap. 4] to (1), written in terms of Φ , with $G = 0$, gives

$$\alpha \nabla^2 \bar{\Phi}(\mathbf{x}) = \bar{\Phi}(\mathbf{x})p - \Phi_0, \quad (2)$$

66 where $\alpha = K/S_s$ is hydraulic diffusivity [L^2/T], p is the Laplace transform parameter [T^{-1}],
 67 $\bar{\Phi}(\mathbf{x})$ is the transformed discharge potential [L^3], and Φ_0 is the initial value of Φ . To render (2)
 68 homogeneous we set $\Phi_0 = 0$; non-zero initial conditions are introduced using impulse area
 69 sources at $t = 0$ [22]. The governing equation in Laplace space is the Yukawa [23] or modified
 70 Helmholtz equation

$$\nabla^2 \bar{\Phi}(\mathbf{x}) - \kappa^2 \bar{\Phi}(\mathbf{x}) = 0, \quad (3)$$

71 where $\kappa^2 = p/\alpha$ [L^{-2}] is analogous to the wave number in wave propagation problems [24, Sect.
 72 1.1.2], or alternatively $\kappa = 1/(Z_0 K)$, where Z_0 is the mechanical analog of impedance [25, Chap.
 73 7].

74 2.1. LAPLACE-SPACE CONVOLUTION

75 Duhamel's theorem [26, Chap. 5] states that temporal behavior of a function can be obtained
 76 from convolution of the impulse response, $\Phi_{\text{imp}}(\mathbf{x}, t)$, and a time behavior, $g(t)$, through the
 77 convolution integral

$$\Phi_{\text{gen}}(\mathbf{x}, t) = \int_0^t \Phi_{\text{imp}}(\mathbf{x}, t - \tau) g(\tau) d\tau. \quad (4)$$

78 LT-AEM elements are derived in Laplace space where (4) becomes

$$\bar{\Phi}_{\text{gen}}(\mathbf{x}, p) = \bar{\Phi}_{\text{imp}}(\mathbf{x}, p) \bar{g}(p). \quad (5)$$

79 When convolution is performed in the time domain [15], each different time behavior (e.g.,
 80 constant, pulse, or linearly-increasing in time) requires approximation of (4). LT-AEM allows for
 81 separate handling of the time (\bar{g}) and space ($\bar{\Phi}_{\text{imp}}$) behavior of elements through the numerical
 82 inverse Laplace transform.

83 Many useful $\bar{g}(p)$ functions can be found tabulated in the literature (e.g., [27, Chap. 29], [21,
 84 App. A], and [26, Sect. 7.3]); piecewise linear or constant functions can be used to describe fairly
 85 general behavior. While it would be possible to perform the convolution in the time domain, we
 86 expect that this increases the required effort. The time-domain convolution integral (4) requires
 87 integrating $0 \leq \tau \leq t$, essentially equivalent to time-marching required by an initial value
 88 problem (e.g., explicit finite differences in time).

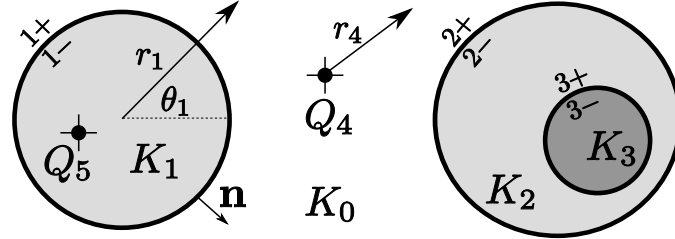


Figure 1. Example with 3 circular elements with different K (background K_0) and 2 prescribed point sources, Q_4 and Q_5 .

89 2.2. BOUNDARY MATCHING

90 Elements are mathematical entities that represent physical objects in the flow system. AEM
 91 and LT-AEM use non-intersecting elements to represent areas of differing properties and source
 92 terms (see Figure 1), by enforcing head and normal flux continuity along the element boundaries.
 93 Each LT-AEM element is derived with implied zero initial condition and zero effect at large
 94 distance. The elements in Figure 1 will be used as an example, 2 point sources of prescribed
 95 strength and 3 circular regions of different K with unknown strengths. Head matching consists
 96 of setting

$$\bar{h}_{\text{tot}}^+(r_{n0}) = \bar{h}_{\text{tot}}^-(r_{n0}); \quad (6)$$

97 the total head, $\bar{h}_{\text{tot}}^\pm = \sum_k \bar{h}_k^\pm$, interior (-) and exterior (+) to the element boundary being set
 98 equal along the boundary of element n , $r_n = r_{n0}$. Head matching along the circumference of
 99 element 2 is expressed in terms of $\bar{\Phi}$ as

$$\frac{1}{K_0} \left[\bar{\Phi}_2^+ + \bar{\Phi}_1^+ + \bar{\Phi}_4 \right]_{r_{20}} = \frac{1}{K_2} \left[\bar{\Phi}_2^- + \bar{\Phi}_3^+ \right]_{r_{20}}, \quad (7)$$

100 where the subscript indicates a local coordinate system, and the super-scripted sign indicates
 101 the side of a two-sided element. Point source Q_5 and the insides of circles 1 and 3 do not appear
 102 in this expression, as they are neither immediately internal nor external to element 2. This con-
 103 vention is used by other AEM applications that solve the modified Helmholtz equation [28, 19]
 104 and is equivalent to the non-overlapping domain decomposition approach called substructuring
 105 [29]; it allows regions with different coefficients in the governing equation to be matched, since
 106 they cannot be combined by superposition.

107 Flux matching applies to the same elements as head matching and consists of setting

$$\mathbf{n}_n \cdot \bar{\mathbf{q}}_{\text{tot}}^+(r_{n0}) = \mathbf{n}_n \cdot \bar{\mathbf{q}}_{\text{tot}}^-(r_{n0}) \quad (8)$$

108 where \mathbf{n}_n is the unit boundary normal for element n and $\bar{\mathbf{q}} = -\nabla \bar{\Phi}$ is the Darcy flux; the total
 109 normal flux, $\mathbf{n} \cdot \bar{\mathbf{q}}_{\text{tot}}^\pm = \sum_k \mathbf{n} \cdot \bar{\mathbf{q}}_k^\pm$, is balanced along the boundary of element n . For element 2
 110 in Figure 1, in terms of $\bar{\Phi}$, this yields

$$\left[\frac{\partial \bar{\Phi}_2^+}{\partial r_2} + \frac{\partial \bar{\Phi}_1^+}{\partial r_1} J_{r_1 r_2} + \frac{\partial \bar{\Phi}_4^+}{\partial \theta_1} J_{\theta_1 r_2} + \frac{\partial \bar{\Phi}_4}{\partial r_4} J_{r_4 r_2} \right]_{r_{20}} = \left[\frac{\partial \bar{\Phi}_2^-}{\partial r_2} + \frac{\partial \bar{\Phi}_3^+}{\partial r_3} J_{r_3 r_2} + \frac{\partial \bar{\Phi}_3^+}{\partial \theta_3} J_{\theta_3, r_2} \right]_{r_{20}}, \quad (9)$$

111 where $J_{\theta_1 r_2} = \frac{\partial \theta_1}{\partial x} \frac{\partial x}{\partial r_2} + \frac{\partial \theta_1}{\partial y} \frac{\partial y}{\partial r_2}$ is a Jacobian relating derivatives in two coordinate systems. Each
 112 of these coordinate derivatives in the Jacobian can be computed explicitly from the geometry of

113 the elements. $\bar{\Phi}$ for each element is defined in terms of a local coordinate system (corresponding
 114 to the one used in separation of variables); differentiation with respect to local coordinates
 115 (e.g., $\partial\bar{\Phi}_2^+/\partial r_2$) leads to more concise expressions than working with a single set of coordinates
 116 everywhere.

117 2.3. SOLUTION FOR COEFFICIENTS

118 The modified Helmholtz equation (3) involves 2 or 3 independent variables (depending on the
 119 dimension, D). The eigenfunction expansion solution is the tensor product of the solutions
 120 obtained through separation of variables [3],

$$\bar{\Phi}_k^\pm(\mathbf{x}) = \prod_{i=1}^D \bar{\Phi}_k^\pm(x_i), \quad (10)$$

121 where $\bar{\Phi}_k^\pm(x_i)$ is a sum of eigenfunctions for the coordinate x_i and element k . The orthogonal
 122 eigenfunctions here are special functions (e.g., Bessel [30, Chap. 6–8] and Mathieu [31, Chap.
 123 9] functions). $\bar{\Phi}$ is expanded in eigenfunctions along element boundaries; the solution is then
 124 computed on or away from the boundaries using the coefficients determined from the boundary
 125 expansion. The second-order ordinary differential equations used here have solutions of the form,

$$\bar{\Phi}_k^\pm(x_i) = \sum_{j=0}^{N-1} a_j^{k\pm} \phi_j(x_i) + b_j^{k\pm} \psi_j(x_i) + R_N^k, \quad (11)$$

126 where $\phi_j(x_i)$ and $\psi_j(x_i)$ are the eigenfunctions associated with the j^{th} eigenvalue and coordinate
 127 x_i ; $a_j^{k\pm}$ and $b_j^{k\pm}$ are free coefficients [L^3] to be determined for the \pm side of element k . The
 128 residual, R_N , arises from truncating the infinite expansion. Upon recombination of the solutions
 129 corresponding to the different coordinate variables (10) products of coefficients are consolidated.

130 Equations (10) and (11) constitute an exact expression for $\bar{\Phi}_k^\pm$, since $R_N^k \rightarrow 0$ as $N \rightarrow \infty$
 131 in a least-squares sense [32] if the eigenfunctions form a complete set. Convergence is at least
 132 $\mathcal{O}(N^{-2})$ for smooth functions with continuous first derivatives [3, Sect. 2.3]. The condition of
 133 smoothness is not overly restrictive for physical problems. In cases where discontinuous functions
 134 must be expanded (e.g., intersecting or touching elements), convergence will be degraded, but
 135 the situation can often be improved with series transformation [33] or smoothing [34, Sect. 49]
 136 techniques.

137 Elements either have specified strength (a_j^k and b_j^k are prescribed) or they have total head or
 138 flux specified in a way which depends on the strength of other elements. LT-AEM requires three
 139 steps to compute head or flux. The first step solves for the coefficients of the eigenfunctions using
 140 boundary collocation, based on a desired arrangement of elements, source terms, and material
 141 properties. The solutions to (11) that arise in the current coordinate system are substituted into
 142 (10) at matching points along the element boundaries to obtain expressions for the coefficients
 143 of the elements. In problems with multiple elements, the coefficients must be either estimated
 144 iteratively (a fixed-point iteration over all elements) or using a direct matrix formulation. The
 145 second step evaluates the solution at the desired \mathbf{x} and p , using the coefficients. Once the
 146 Laplace-space solution and its derivatives are computed, the third step computes the time-
 147 domain solution for head and flux at each location using a numerical inverse Laplace transform
 148 algorithm.

149 2.4. LT-AEM IN RELATION TO AEM

150 Although conceptually LT-AEM is an application of AEM to (3), the implementation is different
 151 in several respects. Since (3) contains p , which is complex, the special functions that satisfy (11)
 152 have complex arguments or parameters. Although some inverse Laplace transform methods only

153 require real p , they are usually less successful at inverting discontinuous time behaviors (e.g., [35,
154 Chap. 9] and [36, Chap. 19]), unless the calculations are performed using very high numerical
155 precision [37].

156 Steady 2D AEM traditionally utilize the complex potential formulation, $\Omega = \Phi + i\Psi$ (where
157 Ψ is a streamfunction), based on the Cauchy-Riemann relations. In LT-AEM both $\bar{\Phi}$ and $\bar{\Psi}$
158 are complex, hence this convention is not applicable, although there are analogous Cauchy-
159 Riemann expressions for the modified Helmholtz equation if Φ and Ψ are real valued [23]. For
160 steady flow Ψ coincides with particle traces, but in transient problems streamlines and pathlines
161 are generally different.

162 For steady 2D AEM, an important distinction is made between elements which have an effect
163 at “infinity” and those which do not (functions of Ω with and without a branch cut) [12, Sect.
164 19]. LT-AEM elements are derived considering that at finite time there is no effect at ∞ , which
165 simplifies derivation and implementation. In the limit as $t \rightarrow \infty$ ($p \rightarrow 0$), these elements would
166 have effects at infinite distance (as $p \rightarrow 0$, (3) becomes the Laplace equation). Therefore, in
167 LT-AEM there are no branch cuts to consider or far-field fixed heads that must be set to obtain
168 a solution, as is required for several common elements in 2D steady-state AEM.

169 Lastly, LT-AEM can readily be modified to handle certain distributed source terms. Analo-
170 gous source terms for $\nabla^2\Phi = 0$ also lead to the Helmholtz equation [7, 19, 20], but would require
171 a significant change in the solution approach. Leakage and transient effects must be dealt with
172 approximately [18] or using area sources [15] in traditional AEM for the Laplace equation, but
173 are readily handled with, or lead to, the Helmholtz equation.

174

3. LT-AEM Elements

175 3.1. TAXONOMY

176 Two-dimensional LT-AEM elements can be categorized with respect to:

- 177 1. boundary condition and whether element coefficients are prescribed (i.e., “given” in AEM
178 literature);
- 179 2. element geometry (e.g., line or area);
- 180 3. changes to source terms or constants in the governing equation (e.g., wave number or initial
181 conditions);
- 182 4. element time behavior (e.g., constant, square wave, or pulse).

183 The free parameters for prescribed elements are independent of other elements in the system
184 (e.g., a well with specified pumping rate). Circles, ellipses, and lines usually define regions of
185 different aquifer parameters, and their coefficients must be determined at run-time. Variable
186 time behavior for any type of element is handled in LT-AEM using Laplace-space convolution.

187 3.2. BOUNDARY CONDITIONS

188 We use boundary condition matching to determine free coefficients; boundary conditions can
189 be Dirichlet, Neumann, or of mixed type. Interface (i.e., matching or continuity) boundary con-
190 ditions are posed along boundaries between regions defined by 2D elements. A mixed boundary
191 condition along the circumference of an element is

$$\xi \nabla \bar{\Phi} \cdot \mathbf{n} + \zeta \bar{\Phi} = \bar{F}(s, p), \quad (12)$$

192 where s is arc length along the boundary. Setting $\xi = 0$, $\zeta = 1$ leaves a Dirichlet boundary con-
193 dition; $\bar{F}_D(s, p) = K\bar{h}_{BC}(s, p)$ is the transformed head along the circumference of the element.

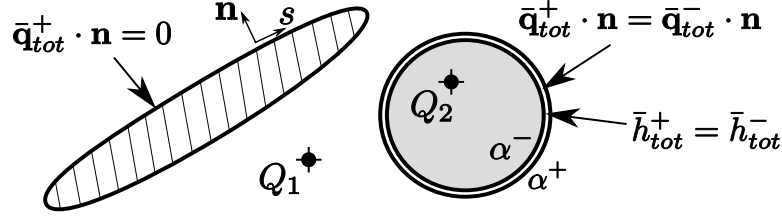


Figure 2. Example with no-flow ellipse, prescribed point sources and circular matching element with different α inside and out (+ and - parts of matching element offset for clarity).

194 With $\zeta = 0$, $\xi = 1$, (12) becomes a Neumann boundary condition; now $\bar{F}_N(s, p) = \bar{q}_{BC}(s, p)$ is
 195 the transformed normal flux on the element boundary.

196 Interface boundary conditions are associated with a two-sided element (see circle in Figure
 197 2); we relate the Neumann and Dirichlet boundary conditions on each side, setting $\bar{F}_N^+(s, p) =$
 198 $\bar{F}_N^-(s, p)$ and $(K^-/K^+)\bar{F}_D^+(s, p) = \bar{F}_D^-(s, p)$, which enforce (6) and (8).

199 We can specify both a Dirichlet and a Neumann condition along the boundary for an elliptic
 200 differential equation because we do not specify a value, but only indicate equality of the inside
 201 and outside. Elements that are not interface conditions have their boundary conditions specified
 202 in terms of total discharge potential (6) or normal flux, (8); if there are at least two elements,
 203 their strengths must be determined simultaneously.

204 To determine element coefficients, M matching points are chosen along the element bound-
 205 aries, creating a system of $2M$ equations (M normal flux $+M$ head), for the $2N$ unknowns.
 206 We use overspecification [38]; by choosing $2M \geq 2N$ the system of equations is solved in a
 207 least-squares sense. Overspecification often produces a smoother solution than $2M = 2N$ does,
 208 and for the same M , N is smaller (i.e., the solution does not require the $2M - 2N$ highest order
 209 basis functions). For these reasons it is utilized in the LT-AEM. We use QR decomposition (as
 210 done in the LAPACK [39] routine ZGELSS) to solve the least-squares problem, rather than
 211 posing the traditional normal equations (e.g., [40, Sect. 5.3] and [41, Chap. 19]).

212 3.3. GEOMETRIC CONSIDERATIONS

213 Table I categorizes elements related to Helmholtz-separable 2D coordinates. Elliptical coordi-
 214 nates are the most general 2D coordinates; polar, parabolic, and Cartesian coordinates can
 215 be obtained by moving the elliptical foci together or moving one or both of the foci to ∞ ,
 216 respectively. In 2D, singular elements are sources or sinks, while areas are defined by finite
 boundaries or infinite lines.

Table I. Helmholtz-separable 2D coordinate systems (e.g., [42], [43, Chap. 1], and [32, Sect. 5.3])

coordinate system	finite boundary	singular element	infinite boundary	modified Helmholtz special functions
Cartesian	<i>none</i>	∞ line	∞ line	exponential
polar (circular)	circle	point	ray	modified Bessel
elliptical	ellipse	line segment	hyperbola	modified Mathieu
parabolic	<i>none</i>	semi- ∞ line	parabola	parabolic cylinder

217

218 3.4. SOURCE TERMS

219 Individual elements or entire domains can be governed by differential equations other than (3);
 220 they can be completely different (e.g., aLplace's equation) or only differ by material properties

221 or the presence of source terms. Source terms can be either homogeneous (functions of $\bar{\Phi}$) or
 222 inhomogeneous (a Poisson term). Homogeneous LT-AEM area sources can be handled without
 223 modification to the solution process, since (3) contains this type of term, additional terms only
 224 change the definition of κ^2 , the wave number. Poisson terms (including Φ_0 in (2)) must be
 225 expressed in terms of a particular solution.

226 3.4.1. Homogeneous leaky aquifer source term

227 Homogeneous source terms arise from effects that are proportional to changes in head or draw-
 228 down in the aquifer. For example, transient leakage from adjacent aquitards, delayed yield in
 229 unconfined systems, and dual-domain behavior all lead to homogeneous source terms [4, Sect.
 230 4.2]. Because the 2D LT-AEM does not represent the third dimension explicitly, Neumann
 231 boundary conditions with respect to the third dimension (e.g., $\partial\bar{\Phi}/\partial z|_{z=0}$) must be represented
 232 as distributed source terms.

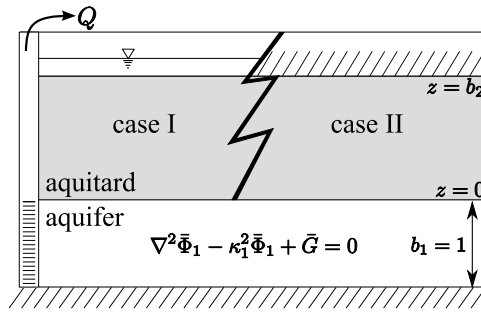


Figure 3. Leaky system diagram

233 Leakage from an adjacent unpumped aquitard leads to a homogeneous distributed source
 234 in 2D. We adapt Hantush's modified leaky system [44] to LT-AEM (see Figure 3). Beginning
 235 with (1) but considering a non-zero source term, G , when expressing the system in terms of $\bar{\Phi}$,
 236 and taking the Laplace transform leads to

$$\nabla^2 \bar{\Phi}_1 - \kappa_1^2 \bar{\Phi}_1 + \bar{G} = 0, \quad (13)$$

237 where subscript 1 indicates the aquifer and 2 the aquitard. Assuming vertical flow in the overly-
 238 ing aquitard (a common assumption when $K_1 \gg K_2$), (3) simplifies to an ordinary differential
 239 equation for $\bar{\Phi}_2$,

$$\frac{d^2 \bar{\Phi}_2}{dz^2} - \kappa_2^2 \bar{\Phi}_2 = 0, \quad (14)$$

240 where the initial value of $\bar{\Phi}_2$ is zero. The head-matching boundary condition at the aquifer-
 241 aquitard interface ($z = 0$) is $\bar{\Phi}_2 = K_2 \bar{\Phi}_1 / K_1$, and at the top the aquitard ($z = b_2$) there is a
 242 no-drawdown condition, $\bar{\Phi}_2 = 0$ (see case I of Figure 3). The solution to (14) that satisfies both
 243 conditions is

$$\bar{\Phi}_2(z) = \frac{K_2 \bar{\Phi}_1}{K_1} [\cosh \kappa_2 z - \coth \kappa_2 b_2 \sinh \kappa_2 z]. \quad (15)$$

244 Differentiating (15) and evaluating it at $z = 0$ gives the vertical flux from the aquitard at the
 245 interface,

$$\bar{G} = \frac{1}{b_1} \left[\frac{\partial \bar{\Phi}_2}{\partial z} \right]_{z=0}; \quad (16)$$

246 when this is substituted into (13), the governing equation in the aquifer becomes

$$\nabla^2 \bar{\Phi}_1 - \left[\kappa_1^2 + \kappa_2 \frac{K_2}{b_1 K_1} \coth \kappa_2 b_2 \right] \bar{\Phi}_1 = 0. \quad (17)$$

247 This can be solved using the same solution techniques used for (3) because the new terms in (17)
 248 are all constants that redefine the wave number. Since the governing equation in the aquitard
 is linear with homogeneous initial and boundary conditions, superposition is valid.

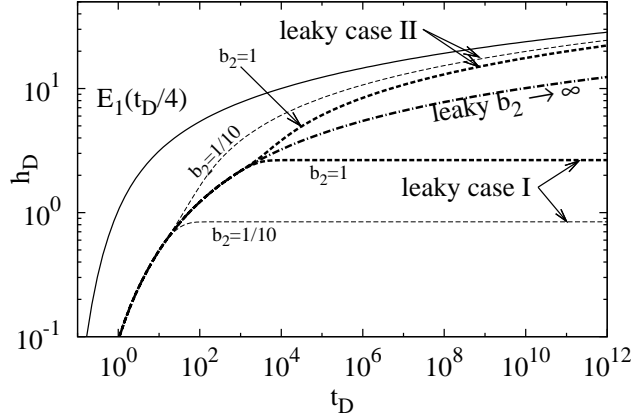


Figure 4. Leaky response at $r = 1$ due to point source (18), comparing results for different aquitard boundary conditions and aquitard thicknesses with the non-leaky E_1 solution; $S_{s2}/S_{s1} = 100$, $K_1/K_2 = 5$.

249 Figures 4, 5, and 6 show effects due to a constant strength finite-radius well source (e.g., [6]
 250 and [45, Sect. 4.3.2], which for (3) is

$$\bar{\Phi}_{\text{well}}(r) = \frac{Q}{2\pi b_1 p} \frac{K_0(r\kappa)}{\kappa r_w K_1(r_w \kappa)}, \quad (18)$$

252 where K_0 and K_1 are modified Bessel functions (see [27, Sect. 9.6] or [30, Sect. 7.2] for properties),
 253 r_w is the pumping well radius [L], and Q is the pumping rate [L^3/T]. to exploit axial symmetry,
 254 plots show dimensionless results; $t_D = tK_1/(S_{s1}r^2)$ is dimensionless time and $s_D = 4\pi|\Phi - \Phi_0|/Q$
 255 is dimensionless drawdown. The curve labeled $E_1(t_D/4)$ represents the non-leaky Theis solution
 256 [46], an exponential integral (see [27, Chap. 5] or [30, Sect. 3.4] for properties). The well solution
 257 (18) to (17) produces the flattening curves in Figure 4.

258 A similar procedure is used to develop a leaky solution with a different aquitard boundary
 259 condition; the equation for a no-flow boundary condition at $z = b_2$ is (case II, the upwardly-
 260 deviating curves in Figure 4)

$$\nabla^2 \bar{\Phi}_1 - \left[\kappa_1^2 + \kappa_2 \frac{K_2}{b_1 K_1} \tanh \kappa_2 b_2 \right] \bar{\Phi}_1 = 0. \quad (19)$$

261 For the thick aquitard case ($b_2 \rightarrow \infty$), $\coth \kappa_2 b_2$ in (17) and $\tanh \kappa_2 b_2$ in (19) simplify to unity
 262 (the middle leaky curve in Figure 4). The effects of the boundary condition at $z = b_2$ are only
 263 observed at later time when the three curves separate (the thin curves in Figure 4 represent an
 264 aquitard 1/10 as thick as the heavy curves — they deviate at an earlier time). The effects of two
 265 aquitards (above and below) can be included, as done by Hantush [44]. The second aquitard
 266 adds another term, analogous to those in (17) and (19).

267 3.4.2. Homogeneous source due to extended form of Darcy's law

268 Higher-order time derivatives in the governing time-domain equation (representing inertia) also
 269 lead to a homogeneous source term in the Laplace-domain. The effect of not considering this
 270 inertia term, in situations where it may be significant (e.g., the coarse gravel-packed region
 271 surrounding a pumping well), may lead to slight over-estimation of storage parameters with
 272 diffusion models. Consider the more complete transient form of Darcy's law (averaged from, or

273 through analogy with, the Navier-Stokes equations (e.g., [47, Sect. 5.10.6] and [48, Sect. 1.5]),
 274 given as

$$\mathbf{q} = - \left(\nabla \Phi + \tau \frac{\partial \mathbf{q}}{\partial t} \right), \quad (20)$$

where τ is the relaxation parameter [T], a property related to the time it takes the system

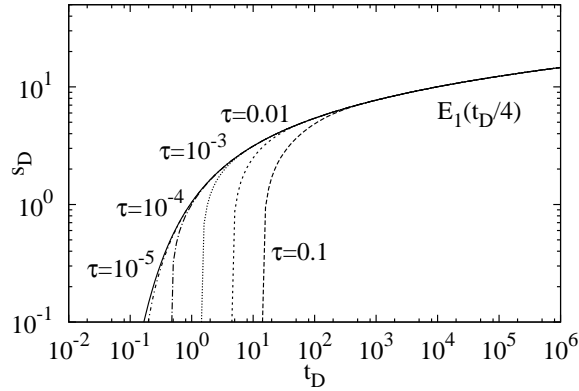


Figure 5. Time-drawdown at $r = 1$ for point source (18) considering inertia effects.

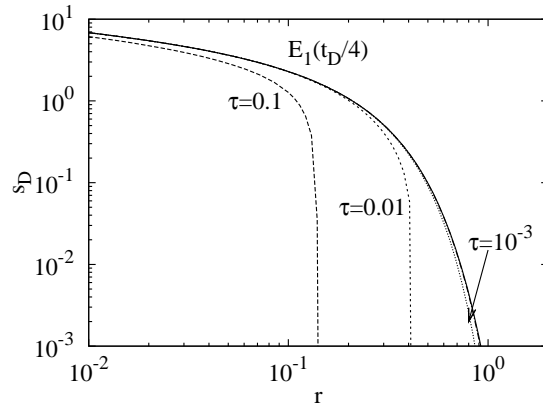


Figure 6. Distance-drawdown at $t = 0.01$ for point source (18) considering inertia effects.

275 to become diffusion-dominated. Typically τ is small and the time derivative term is neglected.
 276 L fqvist and Reh binder [49] define $\tau = K/(ng)$, where n is dimensionless porosity and g is the
 277 acceleration due to gravity [L/T^2]. Combining the Laplace-space mass-conservation equation,
 278

$$-\nabla \cdot \bar{\mathbf{q}} - \kappa^2 \bar{\Phi} = 0, \quad (21)$$

279 with the Laplace transform of (20), the governing equation becomes

$$\nabla \cdot \left[\frac{1}{1 + \tau p} \nabla \bar{\Phi} \right] - \kappa^2 \bar{\Phi} = 0, \quad (22)$$

280 assuming initial head and flux are zero. This can be put into the form

$$\nabla^2 \bar{\Phi} - \left[\kappa^2 + \tau \frac{p^2}{\alpha} \right] \bar{\Phi} = 0, \quad (23)$$

281 which is again similar to (3), but with an additional p^2 term in the wave number. Equation (23)
 282 can be solved for a point source by redefining the wave number in (18) (see Figures 5 and 6).

283 Equation 23 is the transformed damped wave equation, a more general form of the diffusion
 284 equation [25, Chap. 8]. For problems governed by the wave equation, pulses always propagate at
 285 finite speed (e.g., see steep leading edge of s_D surface in Figure 6), while the diffusion equation
 286 allows changes to propagate at infinite speed [50, Sect. 1.2.1]. For example, $s_D = E_1(t_D/4)$
 287 (Theis [46] solution) produces non-zero drawdown at every r for $t > 0$; as $r \rightarrow \infty$ $s_D \rightarrow 0$, so
 288 this discrepancy is usually tolerated. In the damped wave equation τ is inversely proportional
 289 to the maximum propagation velocity squared; as $\tau \rightarrow 0$, the maximum velocity $\rightarrow \infty$, and the
 290 damped wave equation becomes the diffusion equation.

291 The two distributed source terms discussed in this section are illustrated using (18), but
 292 are easily extended to other elements [4, Sect. 4.2]. An elliptical line or area source [22] with
 293 a wave number corresponding to a leaky problem is a trivial extension to the existing line or
 294 area element, thus analytic solutions to other geometries and superpositions thereof are readily
 295 found.

296 3.4.3. Inhomogeneous source terms

297 Area sources can be used to represent constant recharge or discharge, or variable recharge where
 298 the source term is not proportional to aquifer drawdown. For circular elements, Kuhlman and
 299 Neuman [22] showed that $\Phi_0 \neq 0$ can be represented as impulse area sources by decomposing
 300 the solution to the inhomogeneous governing equation into a homogeneous and a particular
 301 solution [12, Sect. 37].

302 The particular solution for an initial condition that is linear or constant in space can be
 303 found by inspection, since the Laplacian of this type of function has zero contribution to the
 304 particular solution. Inhomogeneous terms with more general spatial behavior may be computed
 305 numerically using area integration of the Green's function, through variation of parameters
 306 using known eigenfunctions, or derived as area sources in general functional forms (e.g., 2D
 307 multi-quadric surfaces [51]).

308

4. Elliptical elements

309 Circular LT-AEM elements are given by Furman and Neuman [6], while elliptical elements
 310 are derived here using an analogous procedure [4, Sect. 3.2]. Bakker [28, 52] and Bakker and
 311 Nieber [8] derived elliptical AEM elements for the modified Helmholtz equation. A significant
 312 difference between their elliptical AEM solutions and that given here is the presence of the
 313 complex Laplace parameter, p , which becomes large at small time; this is because p and t
 314 are multiplicative arguments to the exponential in the definition of the Laplace transform (e.g., [34,
 Sect. 4.24] and [26, Chap. 7]).

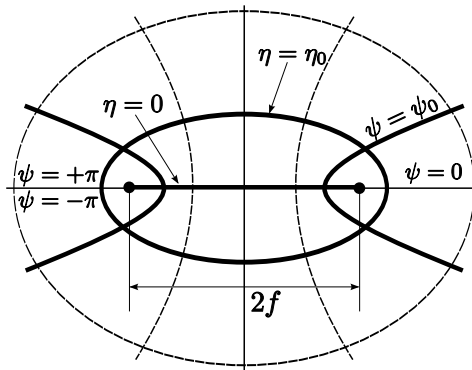


Figure 7. Components of elliptical coordinates (η, ψ) ; f is semi-focal distance.

315

316 Elliptical coordinates (see Figure 7) are defined as $x = f \cosh \eta \cos \psi$ and $y = f \sinh \eta \sin \psi$,
 317 where (η, ψ) are dimensionless elliptical coordinates and f is the semi-focal length $[L]$. The

318 transforms are given succinctly in terms of a conformal map; when $z = x + iy$ and $\zeta = \eta + i\psi$,
 319 the forward transform is $z = f \cosh \zeta$ and the backward transform is $\zeta = \operatorname{arccosh} z/f$. The
 320 multi-valued complex inverse hyperbolic cosine can be expressed as a single-valued function [53]
 321 in the form

$$\zeta = \begin{cases} \ln \left(z/f + \sqrt{(z/f)^2 - 1} \right) & x > 0, \\ \ln \left(z/f - \sqrt{(z/f)^2 - 1} \right) & x \leq 0. \end{cases} \quad (24)$$

322 The modified Helmholtz equation (3) in elliptical coordinates (e.g., [31, Chap. 9], [43, p. 17],
 323 [32, p. 1407]) is

$$\frac{2}{f^2 [\cosh 2\eta - \cos 2\psi]} \left[\frac{\partial^2 \bar{\Phi}}{\partial \eta^2} + \frac{\partial^2 \bar{\Phi}}{\partial \psi^2} \right] - \kappa^2 \bar{\Phi} = 0, \quad (25)$$

324 with the condition that $\bar{\Phi}(\psi) = \bar{\Phi}(\psi + 2\pi)$. Upon substitution of the form $\bar{\Phi}(\eta, \psi) = H(\eta)\Psi(\psi)$,
 325 (25) can be separated into two ordinary differential equations,

$$\frac{d^2 \Psi}{d\psi^2} + (\omega - 2q \cos 2\psi) \Psi = 0, \quad (26a)$$

326

$$\frac{d^2 H}{d\eta^2} - (\omega - 2q \cosh 2\eta) H = 0, \quad (26b)$$

327 where ω is a separation constant (Mathieu characteristic number a or b in Mathieu function
 328 literature) and $q = -f^2 \kappa^2 / 4$ is the Mathieu parameter. These are the angular (26a) and radial
 329 (26b) Mathieu equations. The parameter q is specified through the aquifer properties, element
 330 geometry, and p , while ω is determined to make the solution to (26a) periodic on $\pi \leq \psi < -\pi$.
 331 The special functions that are solutions to (26a) and (26b) are Mathieu functions; see [54, 8, 55]
 332 for characteristic functional plots. Solutions to (25) for $\Re(q) < 0$ are

$$\bar{\Phi}_e^+(\eta, \psi) = \sum_{n=0}^{\infty} a_n \operatorname{Ke}_n(\eta; -q) \operatorname{ce}_n(\psi; -q) + \sum_{n=1}^{\infty} b_n \operatorname{Ko}_n(\eta; -q) \operatorname{se}_n(\psi; -q), \quad (27a)$$

333

$$\bar{\Phi}_e^-(\eta, \psi) = \sum_{n=0}^{\infty} c_n \operatorname{Ie}_n(\eta; -q) \operatorname{ce}_n(\psi; -q) + \sum_{n=1}^{\infty} d_n \operatorname{Io}_n(\eta; -q) \operatorname{se}_n(\psi; -q), \quad (27b)$$

334 where a_n , b_n , c_n and d_n are the coefficients to be determined, Ie_n , Io_n , Ke_n , and Ko_n are the
 335 even (e) and odd (o) radial Mathieu function of first and second kind, and ce_n and se_n are
 336 the even (cosine-elliptic) and odd (sine-elliptic) first-kind angular Mathieu function. Equation
 337 27a only contains the second-kind radial Mathieu functions (Ke_n , Ko_n) which are finite as
 338 $\eta \rightarrow \infty$. Similarly, (27b) only contains the first-kind radial Mathieu function (Ie_n , Io_n) that has
 339 continuous value and slope across the focal line, $\eta = 0$. Because $\Re(q) < 0$ these are modified
 340 Mathieu functions.

341 To simplify the expression for head matching (6) on the boundary of the ellipse, the radial
 342 Mathieu functions are normalized, resulting in

$$\bar{\Phi}_e^+(\eta \geq \eta_0, \psi) \approx \sum_{n=0}^{N-1} a_n \frac{\operatorname{Ke}_n(\eta; -q^+)}{\operatorname{Ke}_n(\eta_0; -q^+)} \operatorname{ce}_n(\psi; -q^+) + \sum_{n=1}^{N-1} b_n \frac{\operatorname{Ko}_n(\eta; -q^+)}{\operatorname{Ko}_n(\eta_0; -q^+)} \operatorname{se}_n(\psi; -q^+), \quad (28a)$$

343

$$\bar{\Phi}_e^-(\eta \leq \eta_0, \psi) \approx \sum_{n=0}^{N-1} c_n \frac{\operatorname{Ie}_n(\eta; -q^-)}{\operatorname{Ie}_n(\eta_0; -q^-)} \operatorname{ce}_n(\psi; -q^-) + \sum_{n=1}^{N-1} d_n \frac{\operatorname{Io}_n(\eta; -q^-)}{\operatorname{Io}_n(\eta_0; -q^-)} \operatorname{se}_n(\psi; -q^-), \quad (28b)$$

344 where the infinite sum has been truncated and the \pm superscripts on q indicate whether it in-
 345 volves aquifer parameters from inside ($-$) or outside ($+$) the ellipse $\eta = \eta_0$. In polar coordinates,

346 a similar set of expressions is derived [16, 6]; they are

$$\bar{\Phi}_c^+(r \geq r_0, \theta) \approx \gamma_0 \frac{K_0(r\kappa^+)}{K_0(r_0\kappa^+)} + \sum_{n=1}^{N-1} \frac{K_n(r\kappa^+)}{K_n(r_0\kappa^+)} [\gamma_n \cos(n\theta) + \delta_n \sin(n\theta)]. \quad (29)$$

347 The first difference between (29) and (28a) is the “even” and “odd” radial functions in elliptical
 348 coordinates. A second difference is the appearance of both an argument (η or ψ) and a parameter,
 349 q^\pm , in (28a) and (28b). Thirdly, both radial and angular Mathieu functions depend on the
 350 coefficients of the partial differential equation (through q), while sine and cosine in (29) do not.

351 4.1. LINE SOURCE

352 An expression for a constant flux line source (along $y = 0$, from $-f \leq x \leq f$) is obtained
 353 from (27a), using only $ce_{2n}(\psi; -q)$ due to symmetry. To simplify flux matching (8) we normalize
 354 by the radial Mathieu function derivative, $Ke'_{2n}(0; -q)$, giving

$$\bar{\Phi}_{\text{line}}(\eta, \psi) = \sum_{n=0}^{\infty} \beta_{2n} ce_{2n}(\psi; -q) \frac{Ke_{2n}(\eta; -q)}{Ke'_{2n}(0; -q)}, \quad (30)$$

355 where β_{2n} are the coefficients to be determined. The boundary condition for a specified flux line
 356 element in elliptical coordinates is

$$\bar{q}_{\text{BC}} = \bar{g}(p) \frac{\bar{\lambda}}{2f} = - \frac{1}{f \sqrt{\frac{1}{2} (\cosh 2\eta - \cos 2\psi)}} \left. \frac{\partial \bar{\Phi}_{\text{line}}}{\partial \eta} \right|_{\eta=\eta_0}, \quad (31)$$

357 where $\bar{\lambda}$ is the transformed constant flowrate [L^3], $2f$ is the length of the line segment, and
 358 \bar{q}_{BC} is the normal flux [L] due to the line source. The metric coefficient in the denominator is
 359 required to preserve the correct dimensions [32, Sect. 1.3]. Differentiating (30) with respect to
 360 η , evaluating it at $\eta = 0$, and using orthogonality over $0 < \psi < \pi$ gives

$$-\bar{g}(p) \frac{\bar{\lambda}}{2} \int_0^\pi \sin \psi ce_{2m}^*(\psi; -q) d\psi = \sum_{n=0}^{\infty} \beta_{2n} \int_0^\pi ce_{2n}(\psi; -q) ce_{2m}^*(\psi; -q) d\psi, \quad (32)$$

361 where $ce_{2m}(\psi; -q)$ has period π and $*$ is complex conjugate. Due to the orthogonality of the
 362 angular Mathieu functions, the integral on the right in (32) is 0 for $m \neq n$ and is defined as $\pi/2$
 363 for $m = n$ [31, Sect. 2.19], reducing the infinite sum to the $2m^{\text{th}}$ term. The expression for the
 364 coefficients is

$$\beta_{2m} = -\bar{g}(p) \frac{\bar{\lambda}}{\pi} \int_0^\pi ce_{2m}^*(\psi; -q) \sin \psi d\psi. \quad (33)$$

365 Expanding ce_{2m}^* in terms of its defining infinite cosine series (A1a), and evaluating the resulting
 366 integral leaves

$$\beta_{2m} = \bar{g}(p) \frac{2\bar{\lambda}}{\pi} (-1)^{m+1} \left[\sum_{r=0}^{\infty} (-1)^r \frac{A_{2r}^{(2m)*}}{1 - (2r)^2} \right], \quad (34)$$

367 where $A_{2r}^{(2m)}$ is a matrix of Mathieu coefficients (see Appendix A). The terms in the infinite sum
 368 quickly become small as r increases and the largest magnitude terms in $A_{2r}^{(2m)}$ occur surrounding
 369 the diagonal $r = m$ (as $q \rightarrow 0$, $A_{2r}^{(2m)}$ becomes a diagonal matrix). Substituting (34) back into
 370 (30) gives the final expression for a constant strength passive line source as

$$\bar{\Phi}_{\text{line}}(\eta, \psi) = \bar{g}(p) \frac{4\bar{\lambda}}{\pi} \sum_{n=0}^{\infty} (-1)^{n+1} \left[\sum_{r=0}^{\infty} (-1)^r \frac{A_{2r}^{(2n)*}}{1 - (2r)^2} \right] ce_{2n}(\psi; -q) \frac{Ke_{2n}(\eta; -q)}{Ke'_{2n}(0; -q)}. \quad (35)$$

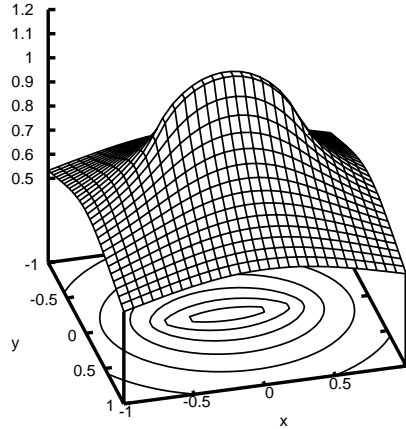


Figure 8. Constant strength line source solution for $f = 0.5$, $Q = 1$, $\alpha = 5.0 \times 10^4$, $t = 0.0125$; h contour interval is 0.1

371 This formulation of the transient line source is valid for any length line source and can take on
 372 different time behaviors through convolution with various $\bar{g}(p)$. Figure 8 illustrates the solution
 373 for $\bar{g}(p) = 1/p$ (a constant strength starting at $t = 0$), using $M = 20$ and an infinite matrix for
 374 the Mathieu functions truncated at 42 terms. Tranter [56] derived a real-valued Mathieu function
 375 solution for the case of an ellipse with a Dirichlet boundary condition. Kucûk and Brigham [57]
 376 applied Tranter’s solution to flow in anisotropic petroleum reservoirs, and Riley [58] derived
 377 expression for flow in a petroleum reservoir to a linear crack. Kuhlman and Warrick [20] derived
 378 a Mathieu function solution for linearized infiltration from an ellipse. Morse and Feshbach [32,
 379 p. 1419–1425] give a solution in terms of Mathieu functions for a Neumann boundary condition
 380 similar to flow through a slot, with real and positive q , while Erricolo [59] shows how these types
 381 of series can be accelerated to minimize the number of Mathieu function evaluations.

382 The flux normal to the line source at $\eta = 0$, illustrated in Figure 8, is compared to the
 383 true boundary condition in Figure 9 for increasing numbers of terms in the Mathieu function
 384 expansion. The numerically-integrated average error, along the boundary of the element, is 0.011
 385 for $N = 4$ and reduces to -3.6×10^{-3} for $N = 12$; the average error decreases slowly beyond
 386 that as more terms are added. The solution converges slowest at the ends of the interval, where
 387 the even function $ce_{2n}(\psi)$ must force the flux to zero.

388 An ellipse or line element expressed in elliptical coordinates using Mathieu functions is useful
 389 as an LT-AEM element with the coefficients of (27a) determined at run time or as an element
 390 for the special case of strength constant in space (35). Using approximate methods, rather than
 391 the appropriate eigenfunctions, may be better suited for intersecting line elements; similar to
 392 those in [28], but using an approximation can accommodate large p values accurately.

393

5. Numerical Inverse Laplace Transform

394 Complex contour integration techniques could be used to analytically compute the time domain
 395 solution from the Mellin contour integral (e.g., [21, Sect. 66] and [45, Sect. 3.2]), defined as

$$\Phi(\mathbf{x}, t) = \mathcal{L}^{-1} \{ \bar{\Phi}(\mathbf{x}, p) \} = \frac{1}{i2\pi} \int_{\sigma_0 - i\infty}^{\sigma_0 + i\infty} \bar{\Phi}(\mathbf{x}, p) e^{pt} dp \quad (36)$$

396 where $\sigma_0 \geq 0$ is larger than the real part of the right-most singularity in $\bar{\Phi}(\mathbf{x}, p)$. Analytic
 397 inversion techniques (e.g., method of residues) are very problem-specific and may only yield
 398 a solution in the form of a slowly converging infinite series; using a numerical \mathcal{L}^{-1} allows

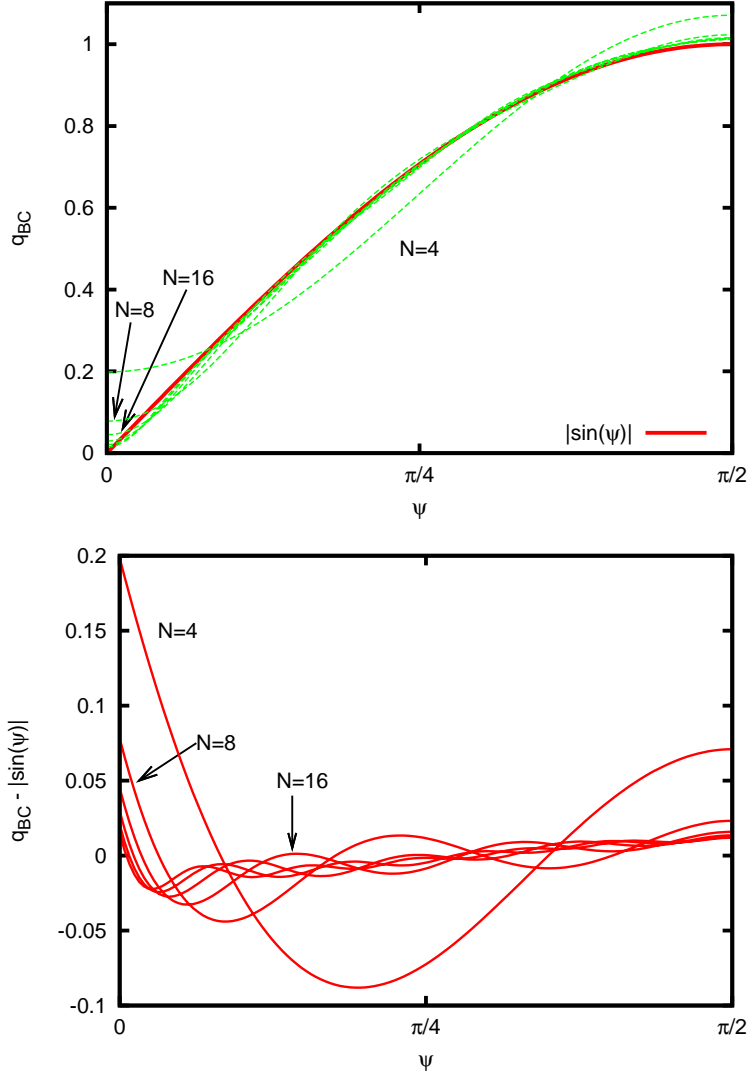


Figure 9. Flux boundary condition and error in Mathieu function expansion for line source in Figure 8.

399 flexibility and generality. See [35, Chap. 9], and the references therein, for general reviews and
 400 comparisons of popular numerical numerical inverse Laplace transform algorithms; Kuhlman
 401 [4, Chap. 5] discusses the details related to several alternative inverse methods in the context
 402 of the LT-AEM.

403 A set of LT-AEM solutions are computed for required values of p , whose optimum values
 404 depend on the algorithm being used. The time-domain solution is then approximated from this
 405 set using a numerical inverse Laplace transform algorithm. Furman and Neuman [6] utilized the
 406 doubly-accelerated Fourier series approach of de Hoog, et al. [60], but no method is universally
 407 best. The Fourier series method can accurately invert an LT-AEM solution over a log-cycle of
 408 t values for a set of p optimized for t_{\max} (the largest t desired). The unaccelerated form of the
 409 Fourier series algorithm is [36, Chap. 19],

$$\Phi(t) \approx \frac{e^{\sigma t}}{T} \sum_{k=0}^{2M'} \Re \left[\bar{\Phi} \left(\sigma + \frac{i\pi k}{T} \right) \exp \left(\frac{i\pi k t}{T} \right) \right], \quad (37)$$

410 where T is a scaling parameter (typically $2t_{\max}$), the first and last terms in the summation are
 411 halved, and σ depends on the locations of the singularities in $\bar{\Phi}$. The argument of $\bar{\Phi}$, the results

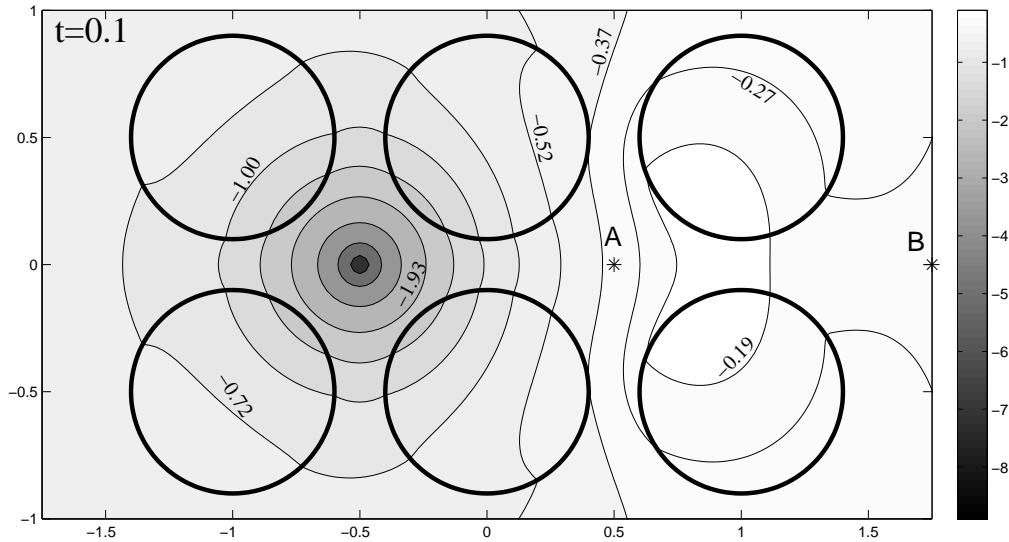
412 of the Laplace-space LT-AEM, are not directly a function of the desired time being inverted,
 413 although the optimal value of T is dependent on t_{\max} . $M \geq 20$ will successfully invert most
 414 time behaviors over a log-cycle of time (e.g., two discontinuities in $g(t)$, representing turning a
 415 pumping well on and off), but smaller M can be used (as low as $M = 3$ for smooth functions)
 416 when T is chosen optimally for each value of time, rather than inverting a whole log-cycle of
 417 times at once.

418 Other algorithms sample the Laplace-space function in different ways. Not all numerical
 419 inverse Laplace transform algorithms are appropriate, depending on problem-specific restrictions
 420 on $\bar{\Phi}$ (e.g., real valued only, or invalid for $\Re(p) < 0$).

421

6. Example: Leaky Circles

422 We simulate six circular regions where κ^2 corresponds to a leaky problem (Case I, (17)); they
 423 are surrounded by material with a wave number $\kappa^2 = p/\alpha$ (Figure 10). A point source at
 424 $(-0.5, 0)$, in the background of the leaky circles, is pumped at a constant rate, starting at $t = 0$.
 425 The circles represent permeable regions in an otherwise impermeable aquiclude separating two
 426 aquifers, the upper aquifer being at constant head. The initial and far field conditions are
 427 $\Phi(r, 0) = \lim_{r \rightarrow \infty} \Phi(r, t) = 0$. Contours of h are logarithmically spaced between -0.01 and
 428 -10 . The pumped aquifer has the same properties everywhere, but the wave number is different
 429 inside the circles, representing the effects of an aquitard with the properties: $K_2 = K_1/2$,
 430 $S_{s2} = (5 \times 10^3)S_{s1}$, $b_2 = 2$. Twelfth-order eigenfunction approximations and 41 values ($M = 20$
 in (37)) of p were used to compute the solution in Figures 10 and 11.



The effects of the leaky circles are clear in the contour plot; the circles behave as area sources, with their recharge rate proportional to the drawdown in the aquifer. The circles “bend” the h contours, reducing the drawdown compared to the non-leaky case.

Figure 10. h contours due to a point source at $(-0.5, 0.0)$ and six case I leaky circular elements (circles represent boundaries of leaky regions) at $t = 0.1$. The aquifer surrounding the circles is non-leaky.

431

432 Figure 11 shows drawdown observed at points A $(0.5, 0)$ and B $(1.75, 0)$ through time,
 433 located in Figure 10. The upper curve represents Theis’ solution [46] (entire domain non-leaky)
 434 and the lower solid curves represent Hantush’s solution [44] (whole domain leaky). The curves

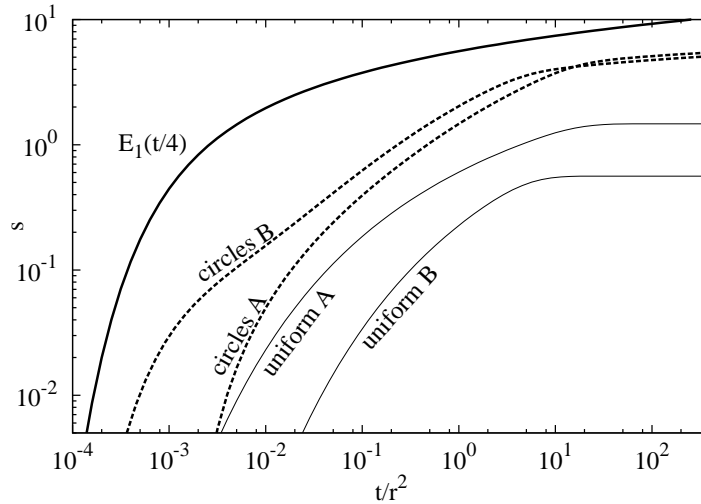


Figure 11. Time-drawdown at two points (see Figure 10 for locations), showing the effects of leaky circles (thick dashed lines) compared to Theis (E_1 , non-leaky) and Hantush (thin solid lines, uniformly leaky) solutions.

435 representing the domain with discontinuous leaky layers (dashed lines) plot between these two
 436 extremes. When the entire domain is leaky, a steady state is reached. With only leaky circles,
 437 the drawdown does not completely flatten out in the figure, the time when the near-steady
 438 portion is reached is shifted compared to Hantush's solutions. For the circles, the approximate
 439 steady-state is approached at a higher value of drawdown, which produces a larger flux from
 440 the aquitard, to compensate for the smaller area producing flux to the aquifer.

441

7. Conclusions

442 AEM and LT-AEM lie between analytic solutions and gridded models in terms of both flex-
 443 ibility and accuracy; they extend some of the elegance of analytic solutions to a broader set
 444 of geometries. LT-AEM additionally utilizes the Laplace transform to achieve flexible analytic
 445 temporal behavior through convolution, while retaining the benefits of AEM.

446 The eigenfunction expansion approach is an elegant method for deriving flexible semi-analytic
 447 solutions for a few geometries. Complex geometries can be approached using approximate tech-
 448 niques borrowed from the boundary and spectral element literature. As examples, the LT-AEM
 449 approach was extended to two problems of potential interest to hydrologists, which are easily
 450 solvable using the same solution techniques used for the standard LT-AEM. These new leaky and
 451 damped-wave solutions exemplify how LT-AEM can be extended to more general aquifer test
 452 analysis scenarios; unconfined, dual porosity, and multi-aquifer flow may be similarly handled
 453 [4, Sect. 4.2].

454 The example illustrates the potential usefulness of LT-AEM for interpreting observed re-
 455 sults from aquifer tests; delivering a flexibility not found in standard analytic aquifer flow
 456 solutions, and an elegance and insight lacking from gridded numerical solutions. Extension of
 457 LT-AEM to three-dimensional flow (i.e., cylindrical and rotational 3D coordinates), elements
 458 with anisotropic material properties, the inclusion of transient particle tracking, the addition of
 459 more aquifer test related elements (e.g., elements with unconfined behavior, wellbore storage,
 460 or a skin layer), and the addition of approximate elements will increase the flexibility and
 461 usefulness of LT-AEM.

462 While the nomenclature and examples used here are specific to hydrogeology, LT-AEM would
 463 be useful for the solution of heat conduction, neutron scattering and other diffusion-dominated
 464 processes. The extension to the damped-wave problem also shows that LT-AEM has the ability

465 to solve additional problems which can be transformed into the modified Helmholtz equation
 466 using the Laplace transform, which includes other non-diffusion processes.

467

Acknowledgments

468 We would like to thank Mark Bakker and two anonymous reviewers for their very detailed and
 469 insightful comments and corrections that greatly improved the quality of the paper. The first
 470 author was supported by the US Geological Survey National Institutes for Water Resources
 471 Grant Program (award 200AZ68G) and the C.W. & Modene Neely fellowship through the
 472 National Water Research Institute.

473

Appendix A: Mathieu Functions

474 To compute Mathieu functions of complex argument, the matrix formulation of the eigenvalue
 475 problem is used here [61–64], solved with LAPACK [39] routine ZGEEV. The traditional con-
 476 tinued fraction approach to solving for the eigenvalues [65] is potentially more efficient than the
 477 matrix method, but it requires an initial guess and is only valid for small Mathieu parameter,
 478 $|q| \leq 4n$, with asymptotic relationships required for larger q [66–68]. The matrix used to compute
 479 the eigenvalues, ω , and eigenvectors, A_r^n and B_r^n , is a truncated infinite matrix, obtained by
 480 substituting the definitions of the angular Mathieu functions back into (26a); the size of the
 481 matrix required is in general proportional to the highest order of Mathieu function needed, the
 482 accuracy desired, and $|q|$ [62, 63].

483 With either the matrix or continued fraction approach, when the Mathieu parameter takes
 484 on complex values, pairs of eigenvalues (and their associated eigenvectors) degenerate at isolated
 485 branch points (double points) in the complex q plane. [65] and [68] discuss the location of and
 486 ramification of these double points.

487 This degeneracy results in the pairs of eigenvectors being less than orthogonal, depending
 488 on the value of q (numerically, the eigenvectors aren't likely to be exactly degenerate). This
 489 behavior is not a problem for the overall convergence of the solution when a more general QR
 490 least-squares solution (e.g., LAPACK routine ZGELSS) is used, which can accommodate this
 491 occasional degeneracy. Numerical inverse Laplace transform methods utilize a set of $\bar{\Phi}(\mathbf{x}, p)$
 492 to compute the time-domain solution. If an entry in this set coincides with a double point of
 493 Mathieu's equation, there will be two non-orthogonal eigenvectors. Because this degeneracy
 494 only affects a pair of the N eigenvectors at one (or possibly two) of the values of p , it is not
 495 critical to the overall performance of the method.

496 Angular Mathieu functions are evaluated from their definitions in terms of infinite sine
 497 and cosine series (second-kind non-period angular Mathieu functions are not useful in our
 498 application), for $\Re(q) < 0$ they are:

$$ce_{2n}(\psi; -q) = (-1)^n \sum_{r=0}^{\infty} (-1)^r A_{2r}^{(2n)} \cos(2r\psi), \quad (\text{A1a})$$

499

$$ce_{2n+1}(\psi; -q) = (-1)^n \sum_{r=0}^{\infty} (-1)^r B_{2r+1}^{(2n+1)} \cos[(2r+1)\psi], \quad (\text{A1b})$$

500

$$se_{2n+1}(\psi; -q) = (-1)^n \sum_{r=0}^{\infty} (-1)^r A_{2r+1}^{(2n+1)} \sin[(2r+1)\psi], \quad (\text{A1c})$$

501

$$se_{2n+2}(\psi; -q) = (-1)^n \sum_{r=0}^{\infty} (-1)^r B_{2r+2}^{(2n+2)} \sin[(2r+2)\psi], \quad (\text{A1d})$$

502 where $A_r^{(n)}$ and $B_r^{(n)}$ are matrices of Mathieu coefficients (both functions of q); each is comprised
 503 of the eigenvectors associated with the n^{th} eigenvalue that provides a periodic solution to angular
 504 Mathieu equation. Even order ($2n$ and $2n + 2$) Mathieu functions are periodic in π , while odd
 505 order Mathieu functions are 2π periodic.

506 Because eigenvectors only define a direction, their lengths must be normalized. An extension
 507 of McLachlan's normalization [31, Sect. 2.21] is used, since it is readily generalized to the
 508 complex case and it produces angular Mathieu function of root mean squared value $1/\sqrt{2}$ over
 509 the entire range of ψ (LAPACK subroutine ZGEEV returns this scaling, additionally scaling the
 510 largest magnitude element of each vector to be real). The Mathieu coefficients are normalized by
 511 $\int_{-\pi}^{\pi} ce_n^*(\psi; -q)ce_n(\psi; -q)d\psi = \int_{-\pi}^{\pi} se_n^*(\psi; -q)se_n(\psi; -q)d\psi = \pi$, where $*$ is complex conjugate;
 512 this makes Mathieu functions degenerate to trigonometric functions as $q \rightarrow 0$.

513 Radial Mathieu functions are best defined in terms of Bessel function product series (conver-
 514 gent for all η). Expressions for them are lengthy, and can be found in the literature [31, Sect.
 515 13.30]. References [27, p. 744] and [55] have tables relating these functions' various names found
 516 in different publications. Derivatives of Mathieu functions are found by applying the derivative
 517 to the definitions; no recurrence relationships exist.

References

1. Stephen R. Kraemer. Analytic element ground water modeling as a research program (1980 to 2006). *Ground Water*, 45(4):402–408, 2007.
2. Otto D. L. Strack and Henk M. Haitjema. Modeling double aquifer flow using a comprehensive potential and distributed singularities 1. solution for homogeneous permeability. *Water Resources Research*, 17(5):1535–1549, 1981.
3. John P. Boyd. *Chebyshev and Fourier Spectral Methods*. Dover Publications, second edition, 2000.
4. Kristopher L. Kuhlman. *Laplace Transform Analytic Element Method*. VDM Verlag, Saarbrücken, Germany, 2008.
5. Charles R. Fitts. Modeling three-dimensional flow about ellipsoidal inhomogeneities with application to flow to a gravel-packed well and flow through lens-shaped inhomogeneities. *Water Resources Research*, 27(5):815–824, 1991.
6. Alex Furman and Shlomo P. Neuman. Laplace-transform analytic element solution of transient flow in porous media. *Advances in Water Resources*, 26(12):1229–1237, 2003.
7. Mark Bakker and Otto D. L. Strack. Analytic elements for multiaquifer flow. *Journal of Hydrology*, 271(1-4):119–129, 2003.
8. Mark Bakker and John L. Nieber. Two-dimensional steady unsaturated flow through embedded elliptical layers. *Water Resources Research*, 40(12):W12406, 2004.
9. R.J. Hunt. Ground water modeling applications using the analytic element method. *Ground Water*, 44(1), 2006.
10. Otto D. L. Strack. Principles of the analytic element method. *Journal of Hydrology*, 226(3-4):128–138, 1999.
11. Otto D. L. Strack. Theory and applications of the analytic element method. *Reviews of Geophysics*, 41(2):1005–1021, 2003.
12. Otto D. L. Strack. *Groundwater Mechanics*. Prentice-Hall, 1989.
13. Henk M. Haitjema. *Analytic Element Modeling of Groundwater Flow*. Academic Press, 1995.
14. Henk M. Haitjema and Otto D. L. Strack. An initial study of thermal energy storage in unconfined aquifers. Technical Report PNL-5818 UC-94e, Pacific Northwest Laboratories, 1985.
15. Willem J. Zaadnoordijk and Otto D. L. Strack. Area sinks in the analytic element method for transient groundwater flow. *Water Resources Research*, 29(12):4121–4129, 1993.
16. James J. Butler and Wenzhi Liu. Pumping tests in nonuniform aquifers: the radially asymmetric case. *Water Resources Research*, 29(2):259–269, 1993.
17. M. Bakker. Transient analytic elements for periodic Dupuit–Forchheimer flow. *Advances in Water Resources*, 27(1):3–12, 2004.
18. Otto D. L. Strack. The development of new analytic elements for transient flow and multiaquifer flow. *Ground Water*, 44(1):91–98, 2006.
19. Mark Bakker and John L. Nieber. Analytic element modeling of cylindrical drains and cylindrical inhomogeneities in steady two-dimensional unsaturated flow. *Vadose Zone Journal*, 3(3):1038–1049, 2004.
20. Kristopher L. Kuhlman and Arthur W. Warrick. Quasilinear infiltration from an elliptical cavity. *Advances in Water Resources*, 31(8):1057–1065, 2008.

21. Ruel V. Churchill. *Operational Mathematics*. McGraw-Hill, third edition, 1972.
22. Kristopher L. Kuhlman and Shlomo P. Neuman. Recent advances in Laplace transform analytic element method (LT-AEM) theory and application to transient groundwater flow. In *Computational Methods in Water Resources*, volume XVI, Copenhagen, Denmark, 2006.
23. R. J. Duffin. Yukawan potential theory. *Journal of Mathematical Analysis and Application*, 35(1):105–130, 1971.
24. Karl F. Graff. *Wave Motion in Elastic Solids*. Dover Publications, 1991.
25. Richard K. Moore. *Wave and Diffusion Analogies*. McGraw-Hill, 1964.
26. Necati M. Özışık. *Heat Conduction*. Wiley-Interscience, second edition, 1993.
27. Milton Abramowitz and Irene A. Stegun, editors. *Handbook of Mathematical Functions with Formulas, Graphs and Mathematical Tables*. Number 55 in Applied Mathematics Series. National Bureau of Standards, 1964.
28. Mark Bakker. Modeling groundwater flow to elliptical lakes and through multi-aquifer elliptical inhomogeneities. *Advances in Water Resources*, 27(5):497–506, 2004.
29. Barry F. Smith, Petter E. Bjorstad, and William Gropp. *Domain Decomposition*. Cambridge University Press, 1996.
30. Larry C. Andrews. *Special Functions of Mathematics for Engineers*. SPIE Press, second edition, 1998.
31. Norman W. McLachlan. *Theory and Application of Mathieu Functions*. Oxford University Press, 1947.
32. Philip McCord Morse and Herman Feshbach. *Methods of Theoretical Physics*, volume 1 and 2. McGraw-Hill, 1953.
33. Cz. Oleksy. A convergence acceleration method for Fourier series. *Computer Physics Communications*, 96(1):17–26, 1996.
34. Cornelius Lanczos. *Applied Analysis*. Prentice-Hall, 1956.
35. Alan M. Cohen. *Numerical Methods for Laplace Transform Inversion*. Springer, 2007.
36. Brian Davies. *Integral Transforms and their Application*. Springer, third edition, 2002.
37. J. Abate and P. P. Valkò. Multi-precision Laplace transform inversion. *International Journal for Numerical Methods in Engineering*, 60(5):979–993, 2003.
38. Igor Janković and R. Barnes. High-order line elements in modeling two-dimensional groundwater flow. *Journal of Hydrology*, 226(3-4):211–223, 1999.
39. E. Anderson, Z. Bai, J. Dongarra, A. Greenbaum, A. McKenney, J. Du Croz, S. Hammarling, J. Demmel, C. Bischof, and D. Sorensen. LAPACK: a portable linear algebra library for high-performance computers. In *Proceedings of the 1990 ACM/IEEE Conference on Supercomputing*, pages 2–11. IEEE Computer Society, 1990.
40. Gene H. Golub and Charles F. van Loan. *Matrix Computations*. Johns Hopkins University Press, third edition, 1996.
41. Charles L. Lawson and Richard J. Hanson. *Solving Least Squares Problems*. SIAM, 1974.
42. F. M. Arscott and A. Darai. Curvilinear co-ordinate systems in which the Helmholtz equation separates. *IMA Journal of Applied Mathematics*, 27(1):33–70, 1981.
43. Parry Moon and Domina E. Spencer. *Field Theory Handbook: Including Coordinate Systems Differential Equations and their Solutions*. Springer-Verlag, 1961.
44. Mahdi S. Hantush. Modification of the theory of leaky aquifers. *Journal of Geophysical Research*, 65(11):3713–3725, 1960.
45. Tien-Chang Lee. *Applied Mathematics in Hydrogeology*. CRC Press, 1999.
46. C. V. Theis. The relation between lowering of the piezometric surface and the rate and duration of discharge of a well using ground-water storage. *Transactions, American Geophysical Union*, 16(2):519–524, 1935.
47. Jacob Bear. *Dynamics of Fluids in Porous Media*. Dover Publications, 1988.
48. Donald A. Nield and Adrian Bejan. *Convection in Porous Media*. Springer, third edition, 2006.
49. Torbjörn Löfqvist and Göran Rehbinder. Transient flow towards a well in an aquifer including the effect of fluid inertia. *Applied Scientific Research*, 51(3):611–623, 1993.
50. Juan Luis Vásquez. *The Porous Medium Equation: Mathematical Theory*. Oxford University Press, 2007.
51. Otto D. L. Strack and Igor Janković. A multi-quadric area-sink for analytic element modeling of groundwater flow. *Journal of Hydrology*, 226(3-4):299–196, 1999.
52. Mark Bakker. Derivation and relative performance of strings of line elements for modeling (un)confined and semi-confined flow. *Advances in Water Resources*, 31(6):906–914, 2008.
53. Raghavendra M. Suribhatla, Mark Bakker, Karl Bandilla, and Igor Janković. Steady two-dimensional groundwater flow through many elliptical inhomogeneities. *Water Resources Research*, 40(4):W04202, 2004.
54. Fayez A. Alhargan. Algorithm 804: subroutines for the computation of Mathieu functions of integer order. *ACM Transactions on Mathematical Software*, 26(3):408–414, 2000.
55. Julio C. Gutiérrez Vega, Ramón M. Rodríguez Dagnino, Antonio M. Meneses Nava, and Sabino Chávez Cerda. Mathieu functions, a visual approach. *American Journal of Physics*, 71(3):233–242, 2003.

56. C. J. Tranter. Heat conduction in the region bounded internally by an elliptical cylinder and an analogous problem in atmospheric diffusion. *Quarterly Journal of Mechanics and Applied Mathematics*, 4(4):461–465, 1951.
57. Fikri Kucuk and William E. Brigham. Transient flow in elliptical systems. *Society of Petroleum Engineers Journal*, 267:401–410, 1979.
58. Michael Francis Riley. *Finite conductivity fractures in elliptical coordinates*. PhD thesis, Stanford University, 1991.
59. Danilo Erricolo. Acceleration of the convergence of series containing Mathieu functions using Shanks transformations. *IEEE Antennas and Wireless Propagation Letters*, 2:58–61, 2003.
60. F. R. de Hoog, J. H. Knight, and A. N. Stokes. An improved method for numerical inversion of Laplace transforms. *SIAM Journal of Statistical Computing*, 3(3):357–366, 1982.
61. L. Chaos-Cador and E. Ley-Koo. Mathieu functions revisited: matrix evaluation and generating functions. *Revista Mexicana de Física*, 48(1):67–75, 2002.
62. Delft Numerical Analysis Group. On the computation of Mathieu functions. *Journal of Engineering Mathematics*, 7(1):39–61, 1973.
63. D. J. Green and S. Michaelson. Series solution of certain Sturm-Liouville eigenvalue problems. *The Computer Journal*, 7(4):322–336, 1965.
64. Jakob J. Stamnes and Bjoern Spjelkavik. New method for computing eigenfunctions (Mathieu functions) for scattering by elliptical cylinders. *Pure and Applied Optics*, 4(3):251–262, 1995.
65. G. Blanch and D. S. Clemm. Mathieu’s equation for complex parameters: tables of characteristic values. Technical report, Aerospace Research Laboratories, US Air Force, 1969.
66. Fayez A. Alhargan. Algorithms for the computation of all Mathieu functions of integer orders. *ACM Transactions on Mathematical Software*, 26(3):390–407, 2000.
67. Felix M. Arscott. *Periodic Differential Equations*. Macmillan, 1964.
68. C. Hunter and B. Guerrieri. The eigenvalues of Mathieu’s equation and their branch points. *Studies in Applied Mathematics*, 64:113–141, 1981.